

THE FACTS ABOUT EL NIÑO

V. M. Ponce and A. V. Shetty



### 1. **DEFINITION**

Variations in climate and weather can significantly affect people's lives. Humans are lin to their ecosystem through climatic events such as extremely cold winters, persistent droughts, or emergencies such as flooding, heat waves, or forest fires. Effective planning requires study and analysis of the climatic variabilities, notably among them, those associated with El Niño.

El Niño is related to the Southern Oscillation, which is a seesaw of air pressures on the eastern and western halves of the Pacific Ocean. Normally the atmosphere above the eas South Pacific is dominated by a persistent high-pressure zone, while a low-pressure zone dominates the western portion. These two systems are coupled; when the pressure rises i the east, it falls in the west, and vice versa. To measure this coupling, meteorologists tak the pressure at Easter Island, about 2,700 miles west of South America, and subtract it fr the pressure at Darwin in northern Australia, and from this difference they calculate the Southern Oscillation Index (SOI).

Under normal conditions, this difference in pressure drives the trade winds from east to v along the equator, i.e., the easterlies. At the same time, high above the ocean surface, the wind circulation is completed as it continues to blow around from west to east. This convection of air is called the Walker Cell circulation, after Sir Gilbert Walker, who identified it in the 1920's. Every 2 to 7 years, the SOI becomes negative. The east end of the pressure seesaw goes down, the west end goes up, and the Walker Cell collapses and sometimes reverses direction. With the collapse of the winds comes the characteristic warm flow of water to east, and the normally cold waters on the South American coast warm up by 2° to 8°C. A of these elements combine to form the phenomenon called El Niño, characterized by a dwindling, and sometimes a reversal, of the trade winds, which causes unusual warming sea-surface temperatures along the tropical eastern Pacific.

### 2. HISTORY

There is a historical record of an El Niño event going back to 1567. Peruvian fishermen along the westernmost shores of South America were the first to give a name to the clim anomaly. In a normal year, the waters were cold and flowed from south to north; howeverevery few years, the waters reversed their flow and warmed up. The phenomenon usually began to peak around Christmas; therefore, the fishermen named the odd weather "El Ni which means "The Child Jesus," in reference to the Christmas holiday.

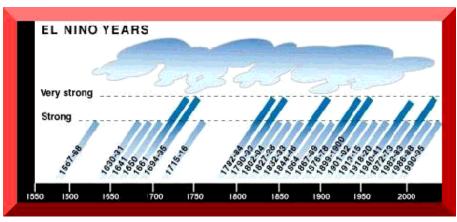
In the early 20<sup>th</sup> century, the concept of El Niño gradually spread through the world's scientific community. At the beginning, scientists believed this strange phenomenon occurred independently of other weather patterns. During the 1920's, while scientists in South America were busy documenting the local effects of El Niño, a British scientist, S Gilbert Walker, separately noted that climate anomalies occurred around the world every few years.

As Sir Walker sorted through world weather records, he recognized some patterns of rain in South America, and associated them with changes in ocean temperatures. He found a connection between barometer readings at stations on the eastern and western sides of th Pacific (Tahiti and Darwin). He noticed that when pressure rises in the east, it usually fal in the west, and vice versa. He coined the term *Southern Oscillation* to emphasize the up and downs of this east-west seesaw effect. He also realized that, under certain barometric conditions, the Asian monsoon seasons were often linked to drought in Australia, Indone India, and parts of Africa, and mild winters in western Canada. He was the first person to claim a connection between the monsoons in India and unusually mild winters in Canada

In the late 1960's, Jacob Bjerknes, a Norwegian meteorologist, proposed that El Niño was the oceanic expression of a large-scale interaction between the ocean and the atmosphere and that the climate anomalies could be better understood as atmospheric *teleconnection* emanating from the warm-water regions of the mid-Pacific along the equator. Starting at about 1975, oceanographers and meteorologists began to combine their efforts to expand and refine the Bjerknes hypothesis by systematically studying El Niño and the Southern Oscillation together in what is now referred to as "El Niño-Southern Oscillation" (ENSC

The 1982-83 El Niño has been widely recognized as the most severe of the century. Nor America experienced wildly unusual weather throughout 1983; Australia experienced massive drought and devastating brushfires; it was one of the worst periods for drought i the sub-Sahelian countries, and the monsoons failed in the Indian Ocean. Total damages

were estimated at between 8 and 13 billion dollars, and 2,000 lives were lost. After the 1982-83 event, there was a minor respite, followed by another event in 1986-88, and yet another in 1990-95, in what appears to be the longest El Niño of record (Fig. 1).



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# Fig. 1 Historical record of El Niño years from the 1500's to 1995.

### 3. CAUSES

The main cause of El Niño involves the full complexity of oceanic-atmospheric interacti on a global scale. While the cause is not fully known, scientists have a good understandi of how it evolves once it has begun. This provides a useful ability to make forecasts six t nine months in advance for some regions.

During normal, or non-El Niño years, the winds at the surface of the Pacific Ocean near equator blow from the northeast and the southeast, converge at the equator, and move we These winds are called *trade winds* because they blow regularly; therefore, trade ships he depended on them to transport their cargo across the ocean. The winds meet in a band parallel to the equator creating a *convergence zone*.

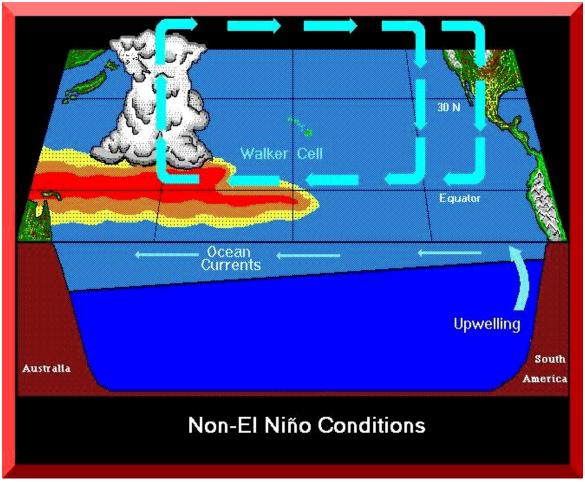
The trade winds drive the ocean currents, causing a flow of surface water in the same direction, from east to west, along the equatorial Pacific. The movement of wind causes warm surface water to build up in the western Pacific. This region of warm surface wate known as the western Pacific *warm pool*. The winds also cause stirring in the upper laye the ocean, which causes warm surface water to be mixed into the deeper ocean. This upp layer, known as the *mixed layer*, is fairly uniform in temperature, salinity and density. TI region which separates the warmer water of the mixed layer from the colder deep ocean water is called the *thermocline*.

The warm-water buildup and the strong surface winds over the western tropical Pacific create a thermocline at approximately 100 m below the surface. In the eastern Pacific, th

cold deep ocean waters are upwelled off the coast of South America to replace the wind-blown warm surface water. This produces a shallow thermocline, approximately at 10-50 m of depth. These upwelled waters are rich in nutrients and produce some of the richer fishing waters in the world.

In the warm pool region of the western Pacific, the warm surface waters and the strong t winds cause large amounts of water to evaporate from the ocean surface. The evaporatio the warm, moist air causes the air to rise and creates an atmospheric low-pressure system the surface. This low-pressure system is known as the *Indonesian Low*. The rising air encounters lower temperature and pressure at higher altitudes, leading to the formation o intense rain-producing cumulus clouds. Further cooling of the rising clouds produces her rainfall over the western Pacific convergence zone. After the water condenses and rains the dry air continues to travel aloft. Some of the air mass travels back towards the east an accumulates over the eastern Pacific. The accumulation of air sinks creating an atmosphehigh-pressure system over the eastern tropical Pacific. Very small amounts of rainfall are experienced in the eastern Pacific because the sinking air prevents the formation of stron cloud systems.

The difference between surface pressures in the eastern and western Pacific produces a *pressure gradient* along the equator. Wind naturally moves across the pressure gradient from high pressure to low pressure. This east-west circulation cell is completed as the air mass continues to move westward along the ocean surface, where it again acquires moist from the evaporating ocean. This pattern of air movement from east to west along the surface and from west to east aloft is the Walker Cell circulation, and it constitutes the east-west component of the large-scale tropical circulation. Figure 2 shows a schematic representation of Non-El Niño (normal) conditions in the tropical Pacific.



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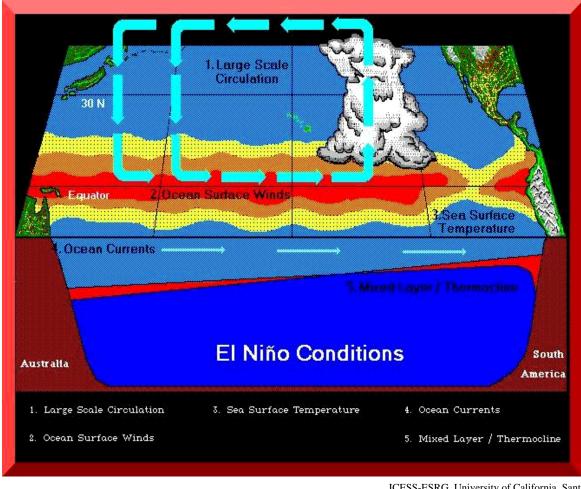
# Fig. 2 During normal years, trade winds blow from east to west, warming the ocean surface along the western Pacific.

The north-south component of the large-scale tropical airflow is the Hadley Cell circulat The Hadley Cell circulation is comprised of two cells which transport heat through the atmosphere from tropical to subtropical latitudes. As the trade winds converge, they crea band of low pressure in the tropics called the *Intertropical Convergence Zone* (ITCZ). Warm moist air rises along the ITCZ carrying with it latent heat. As the air encounters k temperatures at higher altitudes, it cools and the moisture condenses, releasing latent hea The moisture condensation can eventually lead to cloud formation and rainfall. After the moisture condenses and rains out, the dry air continues to travel at high altitudes transporting heat to approximately 30° N and S latitude, where it begins to sink. As the d air sinks, it compresses and heats up due to increased pressure. The air near the surface returns to the equatorial region, gathering moisture along the way and completing the Hadley Cell circulation.

During an El Niño event, several changes occur in the ocean and the atmosphere due to t

dynamic and thermodynamic processes described above. Oceanographers and meteorologists do not yet know what initiates an ENSO event, but they are able to identi and describe the changes associated with ENSO. These changes impact atmospheric circulation patterns and can alter climatic conditions across the globe.

During the early stages of El Niño, the easterlies weaken and retreat eastward, reducing upwelling of cool waters in the eastern Pacific, and allowing the pool of warm water in t west to drift eastward toward South America. As the central and eastern Pacific warms, 1 atmospheric pressure gradient along the equator weakens, and the trade winds diminish ( more. The moist air above the ocean also warms and it becomes buoyant enough to form deep clouds which produce heavy rain along the equator. The change in ocean temperatu causes the major rain zone over the western Pacific to shift eastward. Related adjustment the atmosphere cause barometers to fall over the central and eastern Pacific and rise over Indonesia and Australia, resulting in further weakening and eastward retreat of the easter Figure 3 shows a schematic representation of El Niño conditions in the tropical Pacific.



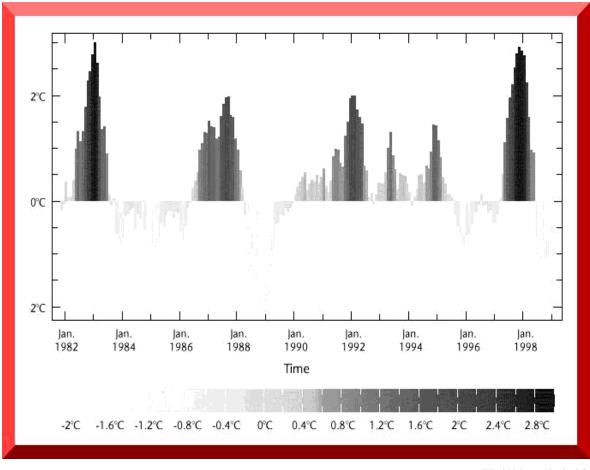
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Fig. 3 During El Niño years, trade winds reverse and blow from west to east, warming the normally cold surface waters of the eastern

#### Pacific.

### 4. PATTERNS

On the average, an El Niño event occurs every 2 to 7 years, but only irregularly, and not predictably as the astronomically controlled tides. The return period of El Niño is varied intensity and duration are also varied. Typically, it lasts anywhere from 14 to 22 months it can be longer or shorter. Every other event tends to be stronger or weaker, with the stre events occurring at 8-15 yr intervals. It decays when there is no longer enough warm wa to sustain the cycle. There are some exceptional cases; a wave of warm water from the 1 El Niño survived for 12 yr, measuring only 8-in high in 1994 and traveling at about 5 mJ Figure 4 shows the 1982-99 time sequence of sea-surface-temperature anomalies in the central equatorial Pacific Ocean.



World Meteorological O1

Fig. 4 1982-99 time sequence of sea-surface-temperature anomalies in the central equ Pacific Ocean.

El Niño often begins early in the year and peaks between the November and following January, but two consecutive events do not behave in the same way. The intervening we and moderate events do not bring disastrous consequences. The event of 1982-83 was unusually strong, equaled only by an event in the late 1800's and the more recent 1997-9 event. Big events like 1982-83 and 1997-98 occur only a few times per century. Evidenc the occurrence of El Niño goes back hundreds of years, in the form of tree-ring analysis, sediment or ice cores, coral-reef samples, and historical accounts. Table 1 shows the El I years of the 20<sup>th</sup> century. There are a total of 30 El Niño years and 26 El Niño events, amounting to a frequency of approximately once every four years.

Table 1. El Niño years of the 20 <sup>th</sup> century.								
1900-01	1902-03	1905-06	1911-12	1914-15	1918-19	1923-24	1925-26	1930-
1939-40	1940-41	1941-42	1946-47	1951-52	1953-54	1957-58	1963-64	1965-
1972-73	1976-77	1977-78	1982-83	1986-87	1987-88	1991-92	1993-94	1995-

K. Amaral, University of Massachu

## 5. IMPACTS

The ocean and atmosphere disturbances created by El Niño can have devastating impact the environment. Changes in the "normal" weather patterns can bring floods to the driest lands on earth, creating lakes where desert plains normally existed. At the same time, El Niño disturbances can bring drought conditions to the wettest regions of the globe, the rainforests, bringing drought stress and sometimes huge forest fires. The effects are glob nature, and extend as far north as Canada and as far south as New Zealand, and over the entire planet from east to west. The effects are felt on many species, from phytoplankton humans. Changes in the normal ocean circulation patterns can cause major disturbances the marine biosphere as sea-surface temperatures change and their associated effects tou all forms of life. The rainfall follows the warm water eastward and results in flooding in Peru and drought in Indonesia and Australia.

The change in weather patterns during an ENSO event alters regions of high and low pressures around the globe. Descending air of atmospheric circulation cells creates high-pressure centers at the surface. The high surface pressures prevent areas of precipitation from moving into this region. When these abnormal high-pressure patterns persist, they lead to drought conditions, depriving the area of rainfall. Droughts generally

occur in the western Pacific during ENSO events, an area normally rich in rainfall. Durin El Niño years, droughts have been observed in many regions, including southeastern Afi India and the northeastern region of South America.

During the 1982-83 ENSO event, Australia was hit particularly hard by drought. After several months without rains, brush fires and dust storms ravaged eastern Australia. Droughts were also endured in parts of Indonesia and the Philippines, where rainforests dried out and subsequently ignited into flames. Summer monsoons from the Indian Ocea where blocked from reaching the Indian subcontinent by a high-pressure system, which brought hardship to the people of Southern India. South Africa and Botswana experience severe drought, with devastating effects on the wildlife and people of the region.

The shifting precipitation patterns associated with El Niño can bring heavy rains to areas accustomed to such amounts. The rains can bring much needed water to dry regions like Peruvian coast. However, the heavy precipitation can create havoc, causing major floodi and sometimes avalanches if the precipitation falls as snow in higher mountainous areas. Some of the devastating impacts of flooding include loss of human lives, destruction of homes, damage to crops and natural vegetation, loss of livestock and the spreading of diseases. In 1982, during the strongest ENSO on record, Peru recorded its greatest rainfa amounts in over 200 years. The rains drenched northern Peru and Ecuador, creating mas floods in the Piura River and in the Quayaquil area. The flood swollen rivers wiped out t prime banana and rice growing regions, swept away vital bridges and flooded towns, overflowing toilet pits and causing virulent diseases to spread thought the affected areas. The rains soaked hillsides and brought thousands of homes tumbling down in the mudsli In Australia, near the end of the 1982-83 ENSO event, and directly after having suffered through a devastating drought, torrential rains swept through and flooded large portions eastern Australia. Tens of thousands of sheep and cattle were stranded in shallow water ( on hilltops surrounded by water. Flooding occurred in many other regions of the world, including Brazil, Europe, China, Cuba and the United States. In the United States much the precipitation fell as snow in the western mountains, delaying the floods until the following spring melt.

Higher sea-surface temperatures in the eastern Pacific provide fertile ground for the development of hurricanes and tropical storms. The eastern Pacific normally has much cooler water, but warm water invades the area during El Niño. Hurricanes are fueled by warm water evaporating off the surface of the ocean. In 1982, Tahiti and its neighboring islands of French Polynesia experienced their first hurricanes in 75 years. Six storms pounded the islands within a five-month period, decimating entire villages on the Tuamc Archipelagos with high winds and waves generated by the storms. These rough seas and increased sea heights combined to make feeding difficult for marine mammals near the shores. Young marine iguanas of the Galapagos Islands were noted by scientists to have smaller sizes than normal. In the same period, 18 storms with wave heights exceeding 6 slammed into the southwestern coast of the United States, damaging over 10,000 homes, uprooting giant kelp beds offshore, and causing extensive coastal erosion which virtually eliminated entire beaches.

Upwelling of deep ocean currents off the coast of Peru create one of the most productive fishing waters in the world. The upwelled ocean currents bring nutrient-rich waters to the

surface. When ENSO conditions prevail, the upwelling is slowed and sometimes even stopped if the event is strong enough. The reduced upwelling decreases the amount of planktonic diatoms which are at the base of the food chain. Without plankton, higher lev life-forms starve or migrate out of the region. Marine species such as zooplankton, maca anchoveta, bait fis, and squid perish from the lack of nutrients, disrupting fishing industr Bird species, which depend on these fish as food sources, also starve or migrate. Often tl leads to nest abandonment by nesting birds. In 1983, up to 85% of the seabirds on the Peruvian coast died. During this same time, the Galapagos and Christmas Islands, home 17 million seabirds, were nearly completely abandoned due to the lack of food. On Christmas Island, this entailed a total reproductive failure for the year.

Sea mammals are also affected by the lack of fish. Sea lions and fur seals have been note have markedly reduced birth rates in the years following an ENSO event. Albatross, boobies, penguins, cormorants and marine iguanas are other species adversely affected b the lack of food. A compounding problem with El Niño is the timing of the events. The warmest sea-surface temperature and its impacts occur when many mammals are breedin At a time when food supplies should be highest, El Niño, diminishes food supplies by reducing upwelling and dampening primary food production. Loss of food has particular detrimental effects on nursing females and their young or unborn. Territorial male mamr suffer due to their inability to move with the migrating food sources.

In spite of the destructive nature of El Niño, some marine creatures do benefit from the climatic disturbances. As a result of the 1982-83 El Niño, scallops accelerated their grow and reached enormous densities. Purple snails and octopuses became more common and shrimp fishery reached its highest level. This is most likely the result of increased run-of from the rivers, which provide a greater abundance of nutrients and decreased predation from a dispersed fish population.

Coral reefs are a notably sensitive ecosystem. The temperature increases create thermal stresses which kill endosymbiotic algae which live in the coral. The algae protects the cc from the sun's damaging ultraviolet (UV) radiation. Without the algae, corals suddenly t white because they are left unprotected against the UV radiation from the sun. The whitening of the coral is known as *coral bleaching*. Within two to four weeks of the bleaching, the coral skeletal structure is destroyed and the plant dies. The 1982-83 El Ni nearly wiped out many eastern Pacific coral communities. Scientist estimated coral mort was close to 98% of the total population. Damage was also observed during this period i reefs in the central and western Pacific, Japan, the Persian Gulf and the tropical western Atlantic, with mortalities reaching 70-90% in some regions. The damage to the coral ree particularly devastating as recovery occurs over time periods of decades to centuries.

### 6. SUMMARY

- El Niño is a global climatic anomaly caused by a weakening of the trade winds alo the tropical Pacific, which leads to warmer-than-normal sea-surface waters along t eastern tropical Pacific.
- The name of El Niño originated with Peruvian fishermen, who noticed that reduce

fisheries were associated with the warmer-than-normal waters. The phenomenon tended to peak around Christmas; thus, they named it "El Niño," for "The Child Jesus."

- The root cause of El Niño is not known with certainty. Atmospheric scientists attri it to the large-scale teleconnections emanating from the warm-water regions of the mid-Pacific along the equator.
- The return period of El Niño is from 2 to 7 years, with an average of 4 years. An eusually lasts from 14 to 22 months. Two consecutive events do not behave in the sway. The strong events occur at 8- to 15-year intervals.
- El Niño can alter the magnitude and frequency of floods and droughts. Typically, i region floods, other regions will experience drought. Floods and droughts can lead extensive damages to life and property. El Niño can also affect biota in diverse ecosystems, including terrestrial and marine.

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